



Carbon dioxide diffuse emission and thermal energy release from hydrothermal systems at Copahue–Caviahue Volcanic Complex (Argentina)



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ABSTRACT

The north-western sector of Caviahue caldera (Argentina), close to the active volcanic system of Copahue, is characterized by the presence of several hydrothermal sites that host numerous fumarolic emissions, anomalous soil diffuse degassing of CO₂ and hot soils. In March 2014, measurements of soil CO₂ fluxes in 5 of these sites (namely, Las Máquinas, Las Maquinitas I, Las Maquinitas II, Anfiteatro, and Termas de Copahue) allowed an estimation that ~165 t of deeply derived CO₂ is daily released. The gas source is likely related to a relatively shallow geothermal reservoir containing a single vapor phase as also suggested by both the geochemical data from the 3 deep wells drilled in the 1980s and gas geoinicators applied to the fumarolic discharges. Gas equilibria within the H–C–O gas system indicate the presence of a large, probably unique, single phase vapor zone at 200–210 °C feeding the hydrothermal manifestations of Las Máquinas, Las Maquinitas I and II and Termas de Copahue. A natural thermal release of 107 MW was computed by using CO₂ as a tracer of the original vapor phase. The magmatic signature of the incondensable fumarolic gases, the wide expanse of the hydrothermal areas and the remarkable high amount of gas and heat released by fluid expulsion seem to be compatible with an active magmatic intrusion beneath this portion of the Caviahue caldera.

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1. Introduction

The poor knowledge of CO₂ fluxes released from natural sources, such as mantle and metamorphic reactions, is one of the most vexing problems in understanding the geological carbon cycle (Berner and Lasaga, 1989). Large uncertainties affect the estimates of global CO₂ flux from volcanoes (Burton et al., 2013 and reference therein) due to the relatively limited flux measurements of volcanic plumes from persistently degassing volcanoes. In addition, the amount of CO₂ not directly related to volcanic craters and released from hydrothermal systems associated with most active volcanic regions is poorly constrained. Recently, an international initiative to fill this gap has been promoted by the scientific community with a project named DECADE (<https://deepcarbon.net/content/deep-carbon-observatory-launches-decade-initiative>), which

supports investigations focused on the study of CO₂ fluxes from active volcanoes. The present study is in the framework of this initiative, being aimed at mapping and quantifying deep-originated CO₂, diffusively discharged from the hydrothermal areas located few kilometers east of the active volcanic system of Copahue (Patagonia, Argentina), where fumarolic discharges and large zones of soil diffuse gas emission occur. A second goal of this study is to provide an estimation of the local geothermal potential.

The development of a quick and reliable technique for the measurements of soil CO₂ fluxes (Chiodini et al., 1998) has recently promoted applications in different fields of geological and environmental sciences. One of the most promising applications of this tool (namely, the accumulation chamber method) regards the use of soil CO₂ flux surveys for geothermal prospecting. This method allows recognizing and characterizing CO₂ flux anomalies at the surface, which are caused by the circulation of hydrothermal fluids at depth. Soil CO₂ fluxes higher than those due to biologic activity are indeed commonly associated with

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the circulation of hydrothermal fluids (Chiodini et al., 1998; Cardellini et al., 2003; Lewicki and Oldenburg, 2005). In addition, recent studies have shown that CO₂ diffuse degassing can provide important and reliable constraints for a correct evaluation of the geothermal potential from hydrothermal areas (Chiodini et al. 2005; Fridriksson et al., 2006; Werner and Cardellini, 2006; Chiodini et al., 2007; Mazot and Taran, 2009; Hernández et al., 2012; Rissmann et al., 2012; Bloomberg et al., 2014; Granieri et al., 2014; Dionis et al., 2015). In particular, the total budget of hydrothermal gases released at the surface can be used for a robust estimation of the minimum amount of geothermal fluids involved at depth in the degassing process. Consequently, the accumulation chamber method represents an effective, rapid and cheap instrumentation for estimating the minimum geothermal potential of an unknown area since the thermal energy naturally transported and released by the fluids can be evaluated.

2. Geological, volcanological and hydrothermal setting

The Copahue–Caviahue Volcanic Complex (hereafter CCVC, 38°S–71°W) is located in the Neuquén Province (Patagonia, Argentina) on a segment of the Andes range, called the South Volcanic Zone (hereafter SVZ: 33.3°–46°S), 30 km east of the main Pleistocene–Holocene volcanic front (Fig. 1). Volcanism in the SVZ is related to the subduction of the Nazca Plate beneath the South American Plate, at rates as high as 10.8 cm y⁻¹ (DeMets et al., 1994; Ramos and Folguera, 2000; Melnick et al., 2006).

The steepening of the oceanic plate subducted in the last 5 Ma resulted in the displacement of the asthenospheric wedge and an asthenospheric upwelling. This process favored a process of crustal thinning that caused the most recent westward migration of the volcanic arc, extensional dynamics and large effusions of basaltic–andesitic magma (Folguera et al., 2006; Yuan et al., 2006).

The CCVC includes the Caviahue Caldera (also known as Caldera del Agrio), a volcano–tectonic depression defined as an intra-arc extensional pull-apart basin (Ramos and Folguera, 2000; Bermúdez et al., 2002; Melnick et al., 2006; Rojas Vera et al., 2010). The pull-apart basin is located at the transition zone between the Liquiñe–Ofqui dextral-slip and the Antifiñir–Copahue fault systems (Lavenu and Cembrano, 1999; Folguera et al., 2004). The former accommodates lateral displacements imposed by the oblique convergence between the Nazca and South American plates from ~46°S to ~38°S (Radic, 2010). The CCVC encompasses the Copahue volcano, a Pleistocene polygenic stratovolcano located in the southwestern rim of the Caviahue Caldera, whose main products are andesites and basalts (Polanco, 2003). The easternmost of the nine NE-oriented summit craters of the Copahue volcano is currently active. During the last 250 years, at least thirteen low-magnitude phreatic and phreatomagmatic eruptions occurred from this crater (Martini et al., 1997; Naranjo and Polanco, 2004). The 1992 and 1995 eruptions mostly consisted of phreatic events characterized by the emission of pyroclastic sulfur. In 2000, a phreatomagmatic eruption, mainly involving juvenile material, occurred (Delpino and Bermúdez, 1993, 2002; GVN, 2000a, 2000b). Since November–December 2011, the discharge rate of fluids from the Copahue active crater increased, whereas sporadic phreatic events have been occurring since July 2012. A major phreatomagmatic–magmatic eruption was observed on December 22, 2012 and a significant degassing is still ongoing (Caselli et al., accepted for publication).

During quiescent periods, the active crater hosts a hot acidic lake (up to 63 °C and pH < 1) (Varekamp et al., 2001; Varekamp et al. 2009; Agosto, 2011; Agosto et al., 2012; 2013). Two acidic hot springs (up to 80 °C and pH = 1–2) discharge in the eastern summit flank of the cone and merge downstream to form the upper Agrio river (pH = 2–3), which flows into the acidified glacial Lake Caviahue (Martini et al., 1997; Gammoms et al. 2005; Varekamp, 2008; Caselli et al., 2005; Agosto, 2011; Agosto and Varekamp 2015).

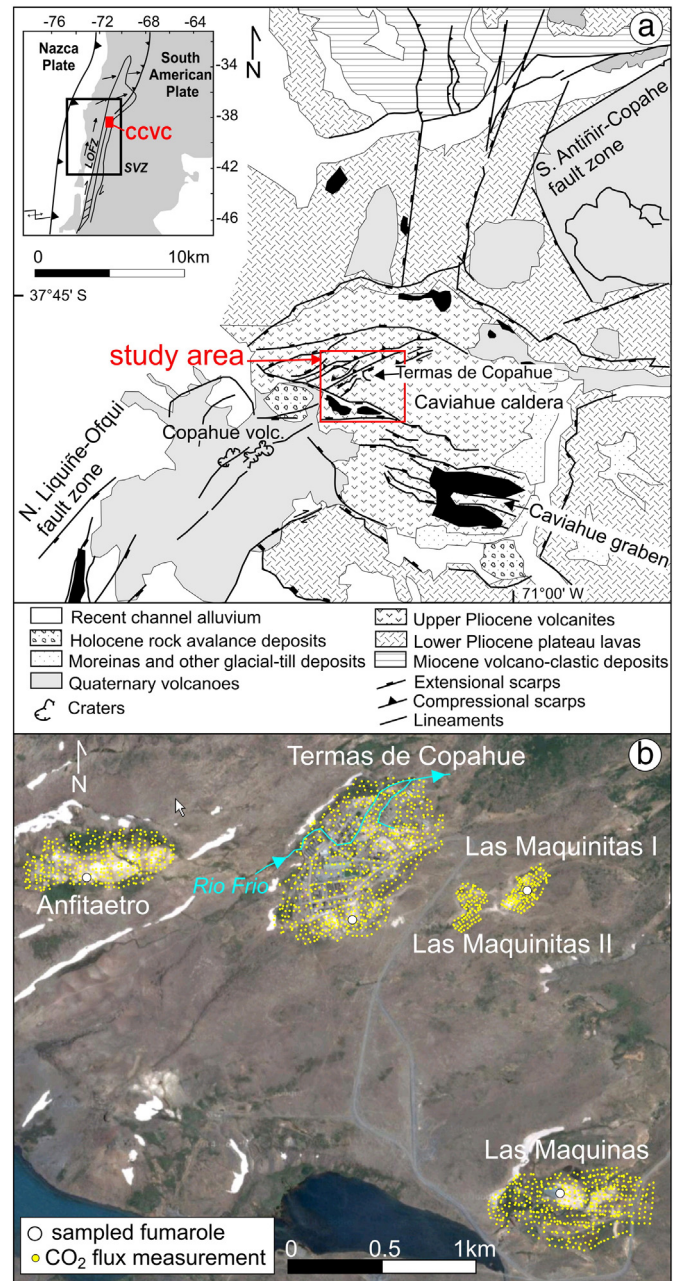


Fig. 1. a) Geological, volcanological and structural setting of the Copahue–Caviahue Volcanic Complex and location of the study area (modified from Folguera et al., 2004); b) location of the surveyed hydrothermal sites.

In March 2014, a remote sensing campaign, carried out by combining MiniDoas and Multigas techniques, revealed the presence of an important gas plume from the crater lake and allowed a rough estimation of the released SO₂ and CO₂, which resulted to be of ~960 and ~640 t d⁻¹, respectively (Tamburello et al., in press).

In the north-eastern flank of the Copahue volcanic edifice, within the Caviahue Caldera, six hydrothermal areas are recognized: Las Máquinas, Las Maquinitas I, Las Maquinitas II, Anfiteatro, Termas de Copahue and Chanco-Co (Mas et al., 1996; 2000; Fig. 1). The hydrothermal activity of some of these sites (Las Máquinas, Las Maquinitas and Termas de Copahue) is so intense that causes a background volcanic tremor as revealed by a seismic array analysis performed in the 2003–2005 period (Ibáñez et al., 2008). Fluids are discharged as boiling, bubbling and mud pools (up to temperatures of 96 °C), fumaroles (up to 130 °C at La Maquinitas I) and large areas of diffuse degassing and hot soils.

According to the recent, comprehensive study of the hydrothermal–volcanic fluids at CCVC by [Agusto et al. \(2013\)](#), the fumarole chemistry suggested that the gas source was associated with boiling processes of a hydrothermal system, mainly fed by meteoric water, although affected by magmatic fluids of mantle signature, as indicated by the relatively high $^3\text{He}/^4\text{He}$ ratios ($R/R_a > 7$).

3. Material and methods

3.1. Sampling and analysis of gas from fumaroles

Fumarolic discharges from the thermal areas of Las Máquinas, Las Maquinillas I and II, Termas de Copahue and Anfiteatro were collected in March 2012 by using pre-evacuated flasks containing a 4 N NaOH solution ([Giggenbach, 1975](#); [Giggenbach and Gouguel, 1989](#)) for the analysis of the major gas species. Vapor condensates and, separately, dry gases were sampled using a condenser, cooled at ~ 20 – 30 °C by cold water. The chemical analyses were carried out at Osservatorio Vesuviano (INGV) laboratories. The gas chemistry of non-absorbed gases, present in the headspace over the NaOH solution, was determined by gas chromatography through a unique injection on two molecular sieve columns (MS 5 Å capillary, $30\text{ m} \times 0.53\text{ mm} \times 50\text{ }\mu\text{m}$; He and Ar as carrier gases) using TCD detectors. Carbon dioxide and sulfur species absorbed in the alkaline solution were analyzed after oxidation via H_2O_2 , by acid–base titration and ion chromatography, respectively (analytical error $\pm 3\%$). Because of reaction in alkaline solution to form COOH^- ([Giggenbach and Matsuo, 1991](#)), CO was analyzed on dry gas samples by gas chromatographic separation with a MS 5 Å $1/8 \times 50$ in column (He as carrier gas) coupled with a high-sensitivity Reduced Gas Detector (HgO). The analytical results are reported in [Table 1](#). In the [Table 1](#) the $^3\text{He}/^4\text{He}$ isotopic ratios, expressed as R/R_a where R is the measured $^3\text{He}/^4\text{He}$ ratio and R_a is that of air (1.39×10^{-6}), are from [Agusto et al. \(2013\)](#).

3.2. Fumarolic flux

A well defined gas plume, suitable for the determination of the CO_2 flux discharged by the fumarolic vent, was found at Las Maquinillas I. Here, the technique proposed by [Aiuppa et al. \(2013\)](#) was applied. The method consists on the measurement of the Integrated Column Amount (ICA, kg m^{-1}) of CO_2 that is subsequently multiplied by the plume transport speed (m/s) to calculate the flux. The concentration in the plume of CO_2 , as well as that of other gases (not discussed here), was measured with a portable MultiGAS system ([Aiuppa et al. 2013](#) and references therein) along the horizontal and vertical axes of an orthogonal cross-section of the plume. During the measurements the plume was sub-horizontal as the wind was blowing to the East with constant speed. We calculated the average CO_2 concentration of ~ 90 samples (0.5 Hz sampling rate) every meter on an 8 m long horizontal axis, and every 0.4 m on a 2.4 m high vertical axis. The gas velocity was determined by tracking the transport speeds of individual gas puffs on a video recorder with a Nikon D90 video camera. The measured plume speed of $6.9 \pm 2.2\text{ m s}^{-1}$ leads to a CO_2 flux of $3.2 \pm 1.1\text{ t d}^{-1}$.

Table 1

Chemical composition of the fumaroles of the surveyed areas (March 2012). Gas concentrations are expressed in $\mu\text{mol/mol}$, helium isotopes as R/R_a ($^3\text{He}/^4\text{He}_{\text{sample}}/^3\text{He}/^4\text{He}_{\text{air}}$). Equilibrium temperatures were calculated within the $\text{H}_2\text{O}-\text{H}_2-\text{CO}_2-\text{CO}-\text{CH}_4$ gas system ($T_{\text{H-C-O}}$; [Tassi et al., accepted for publication](#)), and the geothermometer based on the CO/CO_2 ratio ($T_{\text{CO-CO}_2}$; [Chiodini et al., 2015](#)).

Name	T °C	H ₂ O	CO ₂	H ₂ S	N ₂	CH ₄	H ₂	He	CO	$^3\text{He}/^4\text{He}$	$T_{\text{H-C-O}}$ °C	$T_{\text{CO-CO}_2}$ °C
Las Máquinas	96	973,000	25,200	189	511	540	395	0.32	0.057	7.04	203	206
Las Maquinillas	160	976,000	23,100	214	521	290	312	0.32	0.049	6.97	210	204
Termas de Copahue	95	975,000	24,100	212	449	286	319	0.30	0.052	7.01	210	204
Anfiteatro	92	989,000	9470	213	4045	641	184	0.24	0.068	4.93	258	244

3.3. Soil CO_2 flux and temperatures

Soil CO_2 flux (ϕCO_2) and temperatures (1763 measurements) were measured at Las Máquinas, Las Maquinillas I and II, Anfiteatro, and Termas de Copahue (total investigated area = 1.21 km^2 ; [Fig. 1b](#)). The degassing area of Chanco-Co was not investigated due to logistical problems.

Soil CO_2 fluxes (ϕCO_2) were measured using two accumulation chamber devices developed and calibrated at the laboratories of Osservatorio Vesuviano and University of Perugia. The two equipments, operating in a dynamic mode as described in [Chiodini et al. \(1998\)](#), consist of: 1) a metal cylindrical vessel (the chamber, AC), 2) an Infra-Red (IR) spectrophotometer, 3) an analog-digital (AD) converter, and 4) a palmtop computer. The AC has a volume of $\sim 2.8\text{ L}$ and is equipped with a ring-shaped perforated manifold re-injecting the circulating gas to ensure the mixing of the air in the chamber. The IR spectrometers consist of LICOR Li-800 and LICOR Li-820 detectors equipped with sensors operating in the range 0–20,000 ppm of CO_2 . The soil gas circulates from the chamber to the IR sensor and vice versa by a pump ($\sim 1\text{ L min}^{-1}$). The CO_2 concentration inside the AC is acquired every 250 ms. The signal is converted by the AD and transmitted to a palmtop computer, where a CO_2 concentration vs. time diagram is plotted in a real time. The ϕCO_2 is computed from the rate of CO_2 concentration increase in the chamber (dC_{CO_2}/dt), according to the following equation:

$$\phi\text{CO}_2 = cf \times dC_{\text{CO}_2}/dt. \quad (1)$$

The proportionality factor (cf) between dC_{CO_2}/dt and ϕCO_2 was determined before the survey during laboratory tests. The ϕCO_2 values, typically from 10 to $10,000\text{ g m}^{-2}\text{ d}^{-1}$, were measured on a “synthetic soil” made of dry sand (10 cm thick) placed inside a plastic box with an open top. The cf factor was then computed as the slope of the linear best-fit ϕCO_2 vs. dC_{CO_2}/dt straight line.

Soil temperature was measured at the depth of 10 cm by means of a thermocouple equipped with a metallic probe.

The extension of the five surveyed areas, together with the number and the range of the CO_2 flux (ϕCO_2) measurements for each area, is reported in [Table 2](#). The complete set of the ϕCO_2 data is available in the supplementary material (SM1).

3.4. Soil CO_2 fluxes and temperature data processing

The ϕCO_2 data were used to compute the total CO_2 release from the deep volcanic–hydrothermal source and to map its spatial distribution, as well as that of the soil temperature, by applying the Graphical statistical approach (GSA) and the sequential Gaussian simulation (SGS) methods.

Soil CO_2 flux values in hydrothermal areas are characterized by complex statistical distributions, which generally reflect the coexistence of different CO_2 sources such as biogenic and endogenous ([Cardellini et al., 2003](#)). In a logarithmic probability plot, where a straight line describes one log-normal population, these complex distributions result on a curve with n inflection points, which describes the overlapping of $n + 1$ log-normal populations.

Table 2
Main parameters of the five surveyed areas.

Name	Extension (m ²)	No. of points	Mean (min–max) ϕCO_2 (g m ⁻² d ⁻¹)
Las Máquinas	320,823	495	145 (<0.05–7,270)
Las Maquinitas I	45,842	141	78 (<0.05–2,200)
Las Maquinitas II	32,802	103	272 (<0.05–14,330)
Anfiteatro	26,089	346	105 (<0.05–16,560)
Termas de Copahue	575,748	678	195 (<0.05–9,380)
Total	1,212,585	1763	158 (<0.05–16,560)

The GSA method (Chiodini et al., 1998) was used to both partitioning these distributions into the individual log-normal populations and estimating their proportion (f_i), mean value and standard deviation. The partition was performed according to the graphical procedure proposed by Sinclair (1974). Since the computed statistical parameters of the populations (i.e. mean and standard deviation) refer to the logarithm of values, the mean value of ϕCO_2 (M_i) and the central 90% confidence interval of the mean were estimated by means of a Monte Carlo procedure. The estimated mean flux values were used to compute the CO_2 released from the investigated areas and associated with each population by multiplying M_i by the respective covered surface (S_i), the latter being assumed as a fraction of the total surveyed area (S), which corresponds to the relative proportion of the population (i.e. $S_i = f_i \times S$). The total CO_2 release from the entire area can then be obtained by summing up the contribution of each population (i.e. $\sum f_i \times M_i \times S$). Similarly, the central 90% confidence interval of the mean value was used to calculate the uncertainty of the total CO_2 output estimation of each population.

Although the GSA approach is a useful tool for the interpretation of the diffuse degassing process, the results obtained by this method can be affected by some arbitrary choices, as follows: i) the polymodal log-normal distribution of CO_2 flux values is a convenient model for subsequent partitioning. Nevertheless, the statistical distribution of the CO_2 flux can be more complex than that of a simple log-normal distribution, ii) the partitioning procedure does not imply a unique solution, iii) the spatial distribution of the measured values is not considered by this statistical approach, and iv) the interpretations of the CO_2 flux distribution at the tails, especially for high flux values, can highly be affected by a low number of measured values. As a consequence of the latter “choice”, the estimate of the total CO_2 output can be subjected to remarkable differences.

An alternative and more reliable estimation of the total CO_2 output can be obtained from the CO_2 flux mapping by the Sequential Gaussian Simulation (SGS) algorithm provided by the *sgsim* code (Deutsch and Journel, 1998). According to Cardellini et al. (2003) and Lewicki et al. (2005), SGS yields a realistic representation of the spatial distribution of the CO_2 fluxes reproducing the histogram and variogram of the original data.

The SGS method produces numerous equiprobable and alternative simulations of the spatial distribution of the attribute, i.e. CO_2 flux and temperature in this work. Since the SGS procedure requires a multi-Gaussian distribution, original data were transformed into normal distribution by a normal score transform (Deutsch and Journel, 1998; Cardellini et al., 2003). Experimental variograms of the normal scores were computed and modeled for each data set. The models were used in the SGS procedure to create 200 simulations of the normal scores. The simulated normal scores were then back-transformed into values expressed in original data units, applying the inverse of the normal score transform. The average of the values simulated at each cell of the grid in the 200 simulations was used to draw the maps of soil CO_2 flux and soil temperature. For each simulation the total CO_2 release was computed by summing up the products of the simulated value of each grid cell by the cell surface. The mean and the standard deviation of the 200 values of total CO_2 output were assumed to be the characteristic values of the CO_2 release and of its uncertainty, respectively, for each surveyed area.

4. Results and discussions

4.1. CO_2 soil degassing

The investigated areas were characterized by a wide range of CO_2 flux values, which varied from <0.05 g m⁻² d⁻¹ to $>16,560$ g m⁻² d⁻¹ (Table 2). Each data set is reported in the logarithmic probability plots of Fig. 2. These diagrams show the results of the GSA analysis, which includes i) the partitioned log-normal populations (blue straight lines), ii) their proportion, mean and standard deviation, and iii) the theoretical statistical distribution resulting from the combinations of the individual populations (red dashed curves).

The proportion, mean and standard deviation and the total CO_2 output calculated for each population are reported in Table 3.

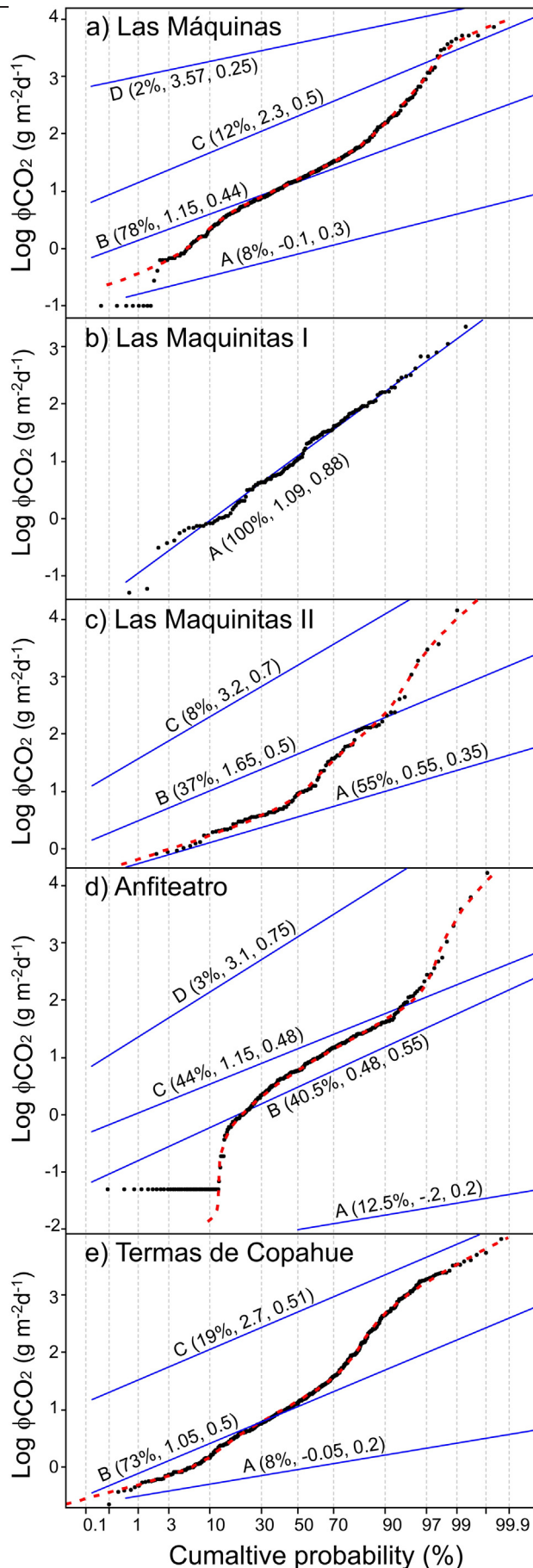
On the basis of the mean flux values characterizing the different populations, an interpretation of the main CO_2 source is reported in Table 3. “Background” refers to CO_2 fluxes related to soil respiration, whereas the term “endogenous” is related to those fluxes fed by volcanic-hydrothermal degassing. The latter includes those populations characterized by high mean ϕCO_2 values, typically in the order of 10^3 g m⁻² d⁻¹, i.e. much higher than those produced by biogenic sources in the soil, which typically are 2–3 order of magnitude lower (e.g., Raich and Schlesinger, 1992; Raich and Tufekcioglu, 2000; Cardellini et al., 2003). At Las Máquinas, Anfiteatro and Termas de Copahue, the distribution of the CO_2 flux values in the probability plots indicates the presence of more than one “background” population (Table 3). The occurrence of different background populations possibly reflects the presence of different soils and vegetation in the surveyed areas. The background populations with the lowest mean values of ϕCO_2 (normally <1 g m⁻² d⁻¹) correspond to fluxes from bare altered soils. Such low values could nevertheless be referred to an endogenous source, although their origin cannot properly be assessed since no isotopic carbon values of the CO_2 efflux (Chiodini et al., 2008) are available. However, it is to be pointed out that contributions by low flux populations to the total CO_2 budget are negligible. The relatively high ϕCO_2 values, which characterize the background populations “B” at Las Máquinas (24 g m⁻² d⁻¹), C at Anfiteatro (26 g m⁻² d⁻¹), and B at Termas de Copahue (22 g m⁻² d⁻¹), are mainly representative of the presence in the surveyed areas of wet soils and peat (Table 3).

The estimated total CO_2 outputs using the GSA approach, i.e. the sum of all contributions from the different populations, range from 4.4 t d⁻¹ (Las Maquinitas I) to 119 t d⁻¹ (Termas de Copahue). The central 90% confidence interval of the mean value is generally large and, especially at Anfiteatro and Las Maquinitas II, it varies one order of magnitude (11 – 110 t d⁻¹ and 4 – 44 t d⁻¹, respectively). These large uncertainties mainly depend on the relatively low number of samples available for the definition of the high-flux populations, which mostly contribute to the total CO_2 output. On the contrary, the computations of the background populations are affected by a lower uncertainty because they are less variable and are defined by numerous samples (Fig. 2, Table 3). Assuming that CO_2 of the background populations is totally derived from shallow biogenic sources (soil respiration, e.g. Raich and Schlesinger, 1992), the total background CO_2 output is of 5.9 t d⁻¹ at Las Máquinas, nil at Las Maquinitas I, 0.09 t d⁻¹ at Las Maquinitas II, 3.3 t d⁻¹ at Anfiteatro and 9.2 t d⁻¹ at Termas de Copahue.

In order to map the CO_2 fluxes and to compute the total gas release using the SGS approach, experimental variograms of the normal scores of the data were computed and modeled for each data set (Table 4). The models were used in the SGS procedure to create 200 simulations of the CO_2 flux according to the computation grids described in Table 4. The obtained CO_2 flux maps are reported in Fig. 3.

All the surveyed areas are characterized by a well-defined diffuse degassing structure (DDS), except at Anfiteatro where the CO_2 fluxes are less spatially organized.

The total SGS-computed CO_2 release ranged from 5 t d⁻¹ (Las Maquinitas I) to 100 t d⁻¹ (Termas de Copahue) with a relatively low



uncertainty ($\leq 10\%$; Table 4). These values can be considered comparable with those obtained by the GSA approach, except for Las Maquinitas II and Anfiteatro, where the SGS estimates are about 50% less than those obtained by GSA. These differences are likely related to an overestimation computed by GSA because a relatively low number of CO_2 flux measurements are available for the definition of the high flux populations. For this reason, the total CO_2 release obtained by the SGS approach was preferred for further computations.

The amount of released endogenous CO_2 (Q_{CO_2}) was computed for each area by subtracting the specific background contribution estimated by GSA to the total CO_2 release estimated by SGS. The computed Q_{CO_2} varies from 5 t d^{-1} (Las Maquinitas I) to 90.8 t d^{-1} (Termas de Copahue) (Table 4).

4.2. Soil temperature distribution

The soil temperature maps obtained by applying the SGS algorithm are reported in Fig. 4 and refer to the temperature at 10 cm depth, concurrently measured with each ϕCO_2 measurement.

Setting aside Anfiteatro, the soil temperature spatial distribution (Fig. 4) in the investigated areas closely mimics that of ϕCO_2 (Fig. 3). A correlation between soil temperature and ϕCO_2 is not surprising because the presence of fumarolic emission favors a massive steam condensation at shallow depth, heating the soil by the latent heat of condensation and causing a flux of incondensable gases (i.e. mostly CO_2) toward the surface (Chiodini et al., 2001; Chiodini et al. 2005). Accordingly, in areas of fumarolic discharges, hot soils and anomalous diffuse soil degassing of incondensable gases, CO_2 flux can be used as a tracer of the whole process allowing an estimation of the total amount of steam and thermal energy involved in the process.

4.3. The hydrothermal system feeding soil diffuse degassing and structural control on DDS

The main fumarolic emissions located in the five surveyed zones (Figs. 1–4) were sampled and analyzed in 2012. The concentration of main and relevant gas species, C and He isotopes and the temperature estimations calculated by gas geothermometry are reported in Table 1. H_2O is by far the main component, being $>97\%$ by volume in all the fumaroles. The second component is CO_2 , followed by minor amount of N_2 , H_2 , CH_4 and H_2S . CO and He concentrations are <1 ppm by volume. The absence of the strong acidic gases (i.e. SO_2 , HCl and HF), which are typical of high temperature fumaroles from active volcanic systems, and the relatively high CH_4 contents suggest that these gases are intimately related to a hydrothermal system. According to Augusto et al. (2013), the fumarolic fluids are originated by boiling of a hydrothermal reservoir, mainly fed by meteoric water. However, the high $^3\text{He}/^4\text{He}$ ratios (R/R_a up to 7.04), the $\delta^{13}\text{C}-\text{CO}_2$ values of $\sim -7\%$ and the N_2/Ar ratios much higher than those of ASW (Air Saturated Water), suggest that He, N_2 and CO_2 are mainly supplied to the hydrothermal system by a magmatic source (Augusto et al., 2013; Tassi et al., accepted for publication). Three deep wells, drilled in the eighties in the frame of a geothermal project (COP-1, COP-2 and COP-3 in Fig. 5; Dellapé and Pando, 1975; Jurío, 1977; Panarello et al., 1988; JICA-EPEN, 1992; Sierra et al. 1992; Mas et al., 2000), provided direct information on the hydrothermal system feeding the CCVC diffuse degassing structures. All the 3 wells, which are located 1–2 km S or W of the studied hydrothermal sites (Fig. 5), reached a deep reservoir of high temperatures (240–260 °C) and a shallower vapor dominated zone at depths of 800–1000 m for which temperatures from 200 to 215 °C were measured and/or estimated with geochemical indicators (Sierra et al., 1990; Panarello, 2002).

Fig. 2. Probability plots of $\text{Log } \phi\text{CO}_2$ for the different hydrothermal sites and partition of the distributions in log-normal populations (blue lines). (For interpretation of the references to color in this figure legend, the reader is referred to the web version of this article.)

Table 3

Estimated parameters and partitioned populations in the 5 surveyed areas.

Name	Population	Proportion (%)	Average ϕCO_2 and 90% confidence interval ($\text{g m}^{-2} \text{d}^{-1}$)	Total diffuse CO_2 output and 90% confidence interval (t d^{-1})
Las Máquinas	A (background)	8	1.01 (0.82–1.23)	0.03 (0.02–0.03)
	B (background)	78	24 (22–27)	5.9 (5.3–6.6)
	C (endogenous)	12	388 (276–538)	15 (11–21)
	D (endogenous)	2	4379 (3058–6038)	28 (20–39)
	Total	100		49 (36–66)
Las Maquinitas I	A (endogenous)	100	95 (49–181)	4.4 (2.2–8.3)
La Maquinitas II	A (background)	55	4.9 (4–6)	0.09 (0.07–.11)
	B (endogenous)	37	86 (56–128)	1 (0.7–1.6)
	C (endogenous)	8	5815 (1298–1,5995)	15.3 (3.4–42)
	Total	100		16 (4–44)
Anfiteatro	A (background)	12.5	<0.1	nd
	B (background)	40.5	6.7 (5.2–8.7)	0.64 (0.49–0.83)
	C (background)	44	26 (21–32)	2.7 (2.2–3.3)
	D (endogenous)	3	5634 (1213–15,022)	40 (9–106)
	Total	100		43 (11–110)
Termas de Copahue	A (background)	8	0.99 (0.89–1.10)	0.05 (0.04–0.05)
	B (background)	73	22 (19–25)	9.2 (8.1–10.4)
	C (endogenous)	19	1000 (784–1270)	109 (86–139)
	Total	100		119 (94–174)

Geothermometric calculations in the $\text{H}_2\text{O}-\text{CO}_2-\text{CH}_4-\text{CO}-\text{H}_2$ gas system (Tassi et al., accepted for publication) indicated that the fumarolic fluids discharged at Piedra Copahue, Las Máquinas and Las Maquinitas equilibrate in a single vapor phase, as actually observed by the geothermal wells, at a temperature of 203–210 °C ($T_{\text{H-C-O}}$ in Table 1). Other computations, based on the CO/CO_2 ratios by applying the same method described in Chiodini et al. (2015), produced similar temperatures ($T_{\text{CO-CO}_2} \sim 204\text{--}206$ °C; Table 1). These estimations are in good agreement with the temperatures measured in the geothermal wells, suggesting the occurrence of a large, probably unique, vapor zone reached by the wells and feeding the hydrothermal manifestations of Termas de Copahue, Las Máquinas and Las Maquinitas (“hot area with evidences of a single phase vapor zone at depth” in Fig. 5). This would explain the remarkable chemical and isotopic homogeneity of the fumaroles from the 3 different sites (Table 1), which are distant a few kilometers from each other (Fig. 5). The three fumaroles show, for example, a similar $\text{H}_2\text{O}/\text{CO}_2$ molar ratio of ~ 40 and $^3\text{He}/^4\text{He}$ of ~ 7 R/Ra.

In order to better understand the role of this vapor zone in the hydrothermal circulation, the structural setting of the zone needs to be considered. The caldera is locally characterized by three fault systems, which are NE–SW, WNW–ESE and NW–SE oriented (Melnick et al., 2006; Rojas Vera et al., 2010; Latinoconsult, 1981; JICA–EPEN, 1992). These three fault systems are arranged in such a way that they constitute the borders of a triangle-shaped horst structure which, according to gravity and electrical resistivity surveys, represents a high conductivity zone of hot fluids circulation (JICA–EPEN, 1992). The geometries of the DDS’s, as defined by the ϕCO_2 distribution, are roughly consistent with these three directions (Fig. 5). In particular at Termas de Copahue, Las Maquinitas I and II the high CO_2 fluxes seem to be mainly distributed along the NE–SW-aligned structures, which correspond to either known faults or faults inferred by this investigation on the basis of diffuse degassing processes active in this area. At Las Máquinas, the DDS

develops along both NE–SW and WNW–ESE structural trends. The general correspondence between the structural trends and the DDS geometries suggests that the emission of the hydrothermal fluids is favored by the fault systems, which cut through the vapor zone, causing the transfer of the deep fluids toward the surface. In Fig. 5, the extension of such vapor zone was roughly delimited: the studied DDS would be located in the northern and eastern limits of this “hot area” with the exception of Anfiteatro which, according to this hypothesis, would be positioned externally with respect to the “hot area”. This is supported by the chemical and isotopic composition of the fumarolic fluids discharged at Anfiteatro, as they significantly differ from the other areas. The Anfiteatro fumaroles are indeed richer in water ($\text{H}_2\text{O}/\text{CO}_2$ molar ratio of ~ 100) and the $^3\text{He}/^4\text{He}$ ratio is significantly lower ($R/Ra \sim 4.9$) than those measured at Las Máquinas, Las Maquinitas and Piedra Copahue.

4.4. Estimation of the thermal energy release

At the Copahue hydrothermal sites, the thermal energy release was estimated by using an approach similar to that described in Chiodini et al. (2001; 2005). The computation was based on (i) the estimation of the pristine $\text{H}_2\text{O}/\text{CO}_2$ ratio ($R_{\text{H}_2\text{O}-\text{CO}_2}$ by weight) of the fluid feeding the soil diffuse gas emission before steam condensation and (ii) the computation of the total steam involved in the process Q_{steam} by multiplying $Q_{\text{CO}_2,d}$ by $R_{\text{H}_2\text{O}-\text{CO}_2}$. In the case of Las Maquinitas, the measured fumarolic CO_2 flux (3.2 t d^{-1}) was added to the diffuse CO_2 output. In each hydrothermal site, Q_{steam} was computed with the reasonable assumption that $R_{\text{H}_2\text{O}-\text{CO}_2}$ is equal to the $\text{H}_2\text{O}/\text{CO}_2$ ratio measured in the fumaroles of the correspondent degassing structure (Table 5). The total amount of steam from each area (Q_{steam}) varies from 285 t d^{-1} at La Maquinitas to 1506 t d^{-1} at Termas de Copahue (Table 5). The

Table 4Relevant parameters of SGS application and estimation of the total CO_2 output from Copahue hydrothermal sites.

Site name	Variogram model, nugget, range (m)	Grid parameters: n. cells, lag (m)	Total CO_2 release \pm standard deviation (t d^{-1})	Endogenous CO_2 release, $Q_{\text{CO}_2}(\text{t d}^{-1})$
Las Máquinas	Spherical, 0.47, 145	35,647, 3	42.7 ± 4.33	36.8
Las Maquinitas I	Spherical, 0.51, 50	45,842, 1	5.01 ± 0.81	5.0(1)
Las Maquinitas II	Spherical, 0.46, 66	32,802, 1	8.30 ± 1.56	8.3
Anfiteatro	Spherical, 0.71, 80	29,748, 3	24.0 ± 2.45	21.7
Termas de Copahue	Spherical, 0.59, 194	63,972, 3	100 ± 5.42	90.8
Total	–	–	180	162.6

(1) At Las Maquinitas I, an additional CO_2 flux of $\sim 3.2 \text{ t d}^{-1}$ was measured from the main fumarolic vent.

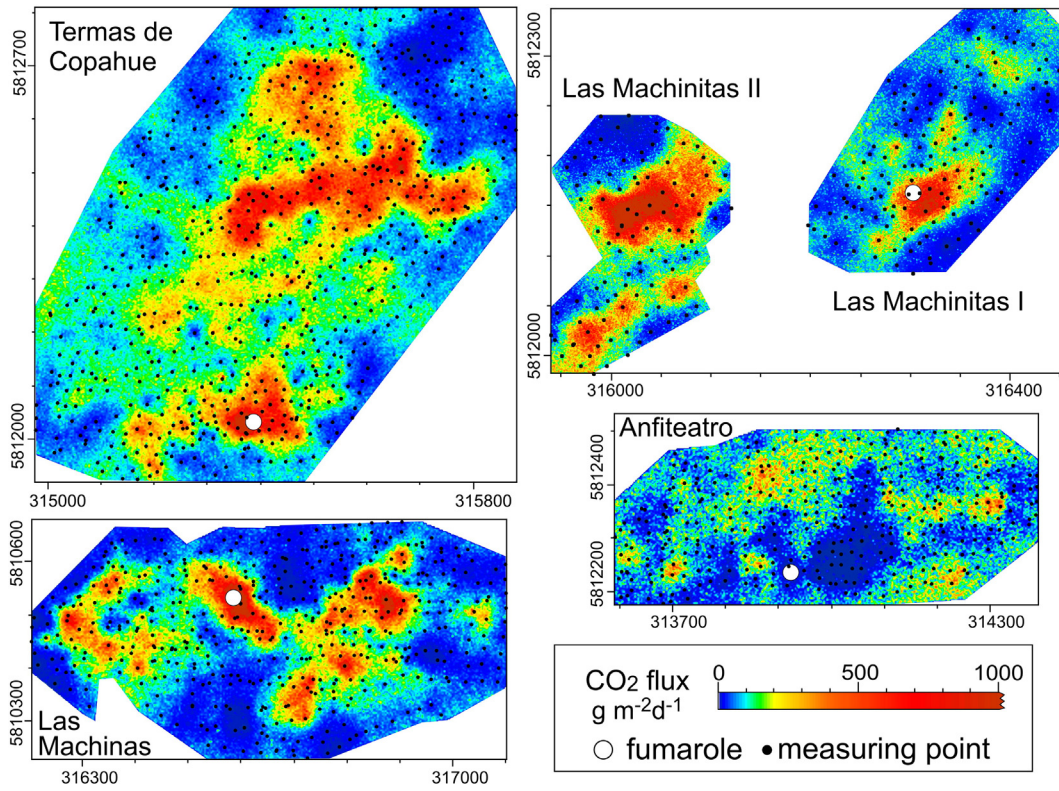


Fig. 3. Maps of the CO₂ flux for the different hydrothermal sites (map coordinates are expressed in m, UTM-WGS84 19S).

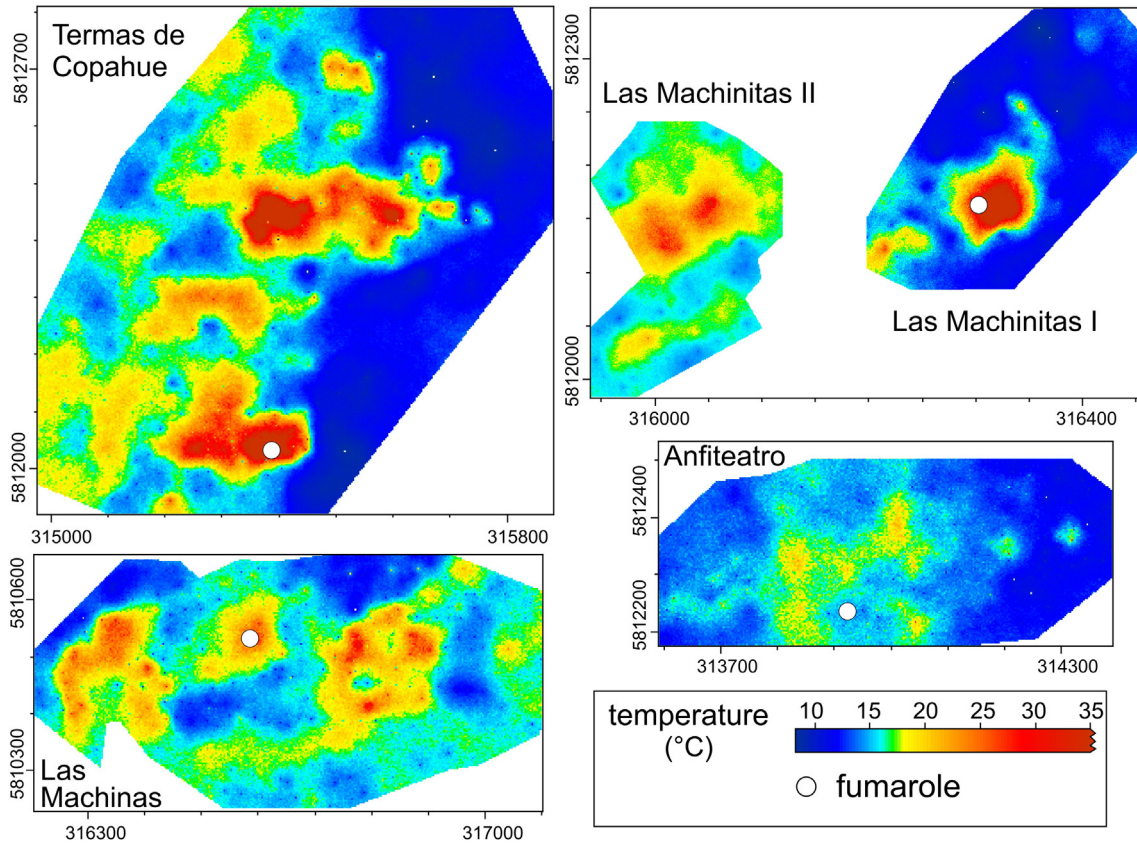


Fig. 4. Maps of soil temperature for the different hydrothermal sites (map coordinates are expressed, in m UTM-WGS84 19S).

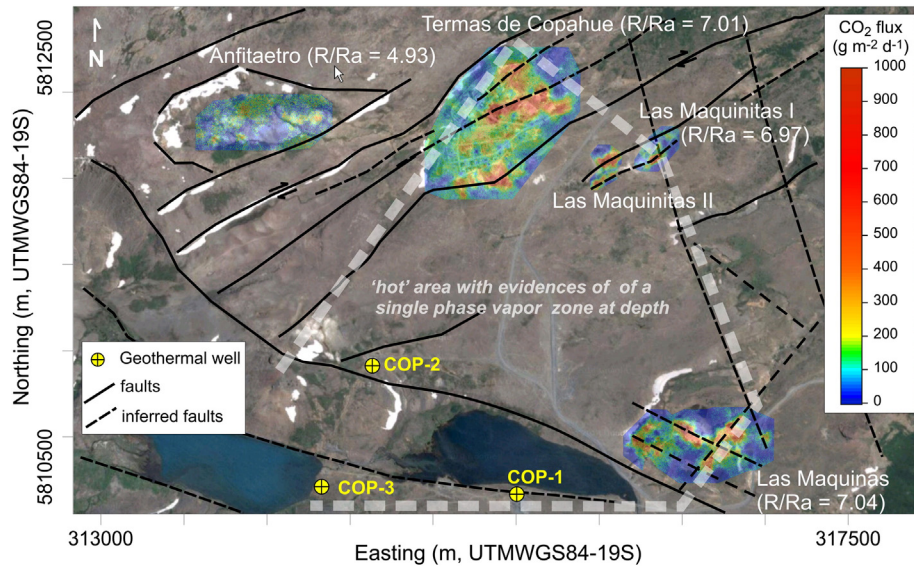


Fig. 5. Structural setting for area compared with the location of DDSs and geothermal wells. The area where we infer the presence at depth of a single phase vapor zone is highlighted.

total thermal release QH_{tot} (Table 5) was calculated by adding three contributions:

- (1) QH_{res} represents the heat released by the $\text{H}_2\text{O}-\text{CO}_2$ gas mixture moving from the reservoir conditions to the condensation zone. The reservoir temperatures were considered equal to 210 °C, whilst the reservoir pressure was assumed that of the saturated vapor at which P_{CO_2} , computed by multiplying $P_{\text{H}_2\text{O}}$ by the measured fumarolic $\text{CO}_2/\text{H}_2\text{O}$ molar ratio, was added. QH_{res} was calculated by multiplying Q_{steam} by the enthalpy difference between the vapor at reservoir conditions and at condensation conditions (0.096 Mpa and 98 °C). The computation was performed using MUFITS software, which allows prediction of $\text{CO}_2-\text{H}_2\text{O}$ mixture properties in a wide range of pressures and temperatures (Afanasyev, 2013). QH_{res} ranges from 0.5 MW to 3.6 MW, thus it is the minor term of the energetic balance of the diffuse degassing structures (Table 5);
- (2) QH_{cond} corresponds to the heat released during steam condensation at subsurface conditions. The QH_{cond} values, which were computed by multiplying Q_{steam} by the latent heat of condensation at 98 °C (2262 J g⁻¹), range from 7.5 MW at Las Maquinitas I and II areas to 39.4 MW at Termas de Copahue (Table 5). QH_{cond} is the main term of the energy budget;
- (3) QH_{cooling} is the heat released as the condensates cool down to ambient temperature. It was estimated by multiplying Q_{steam} by the enthalpy difference between the liquid at 98 °C (enthalpy = 411 J g⁻¹) and at 10 °C (enthalpy = 42 J g⁻¹).

The QH_{cooling} values, from 1.2 to 6.4 MW, were intermediate between those of QH_{res} and QH_{cond} (Table 5).

The total thermal energy release from the five-surveyed zones is 107.5 MW. The highest thermal energy release (49.1 MW) was estimated at Termas de Copahue, where the computed Q_{steam} was 1506 t/d. Here, the production of large amount of condensates is shown by the mass balance calculated for the small creek (Rio Frio, Fig. 1) that enters the village with a flow rate of 560 t d⁻¹ (pH = 6.06, T = 16.2 °C), collects the great majority of the condensation waters and flows out at the rate of 1460 t d⁻¹ (pH = 3.4, T = 22.5 °C). The measured flow rate increment, which is about 60% of the estimated condensate production in the area, appears to be realistically supporting the reliability of our estimation. It is indeed reasonable that part of the condensates is feeding the local aquifer (groundwater circulation).

5. Conclusions

The north-western sector of the Cavihue caldera is characterized by fumarolic emissions associated with zones of anomalously high soil CO_2 diffuse degassing and soil temperature. Five of these sites were investigated and a total discharge of deeply-originated CO_2 of ~165 t d⁻¹ from soil diffuse degassing processes was estimated. The gas source for Termas de Copahue, Las Máquinas, Las Maquinitas I and II is a 800–1000 m deep vapor zone with a temperature of ~200–215 °C, as indicated by both the data of three deep wells drilled in the eighties SW of the natural degassing sites, and gas geothermometry. The occurrence of a unique gas zone feeding the manifestations of the area explains the remarkable compositional homogeneity of the fumaroles, with the exception of those discharging at Anfitaetro, where significant compositional and isotopic differences with respect to the other sites were observed. Using CO_2 as a tracer of the original vapor phase, a natural thermal release of ~77 MW from Termas de Copahue, Las Máquinas, Las Maquinitas I and II was computed, and increases up to ~107 MW when the Anfitaetro degassing zone is considered.

The clear magmatic signature of the incondensable fumarolic gases, the wide expanse of the hydrothermal zones and the remarkable amount of gas and heat released by fluid expulsion, appear to be compatible with an active magmatic intrusion in the subsurface of this portion of the Cavihue caldera. This model well agrees with the proved

Table 5
Heat flux estimation.

Hydrothermal site	$R_{\text{H}_2\text{O}-\text{CO}_2}$	Q_{steam} (t d ⁻¹)	QH_{res} (MW)	QH_{cond} (MW)	QH_{cooling} (MW)	QH_{tot} (MW)
Las Máquinas	15.8	581	1.2	15.2	2.5	18.9
Las Maquinitas I, II ^a	17.3	285	0.5	7.5	1.2	9.1
Termas de Copahue	16.6	1506	3.2	39.4	6.4	49.1
Anfitaetro	42.7	927	2.2	24.3	4.0	30.4
Total	–	3244				107.5

^a At Las Maquinitas Q_{steam} includes the contribution of the main fumarolic vent which was computed in 55 t d⁻¹ by multiplying the measured CO_2 flux (3.2 t d⁻¹) by $R_{\text{H}_2\text{O}-\text{CO}_2}$.

occurrence of volcanic seismic tremor associated with the hydrothermal systems of the Copahue–Caviahue Volcanic Complex (Ibáñez et al., 2008; Forte et al., 2012).

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